## TES observations of tropospheric ozone as a greenhouse gas

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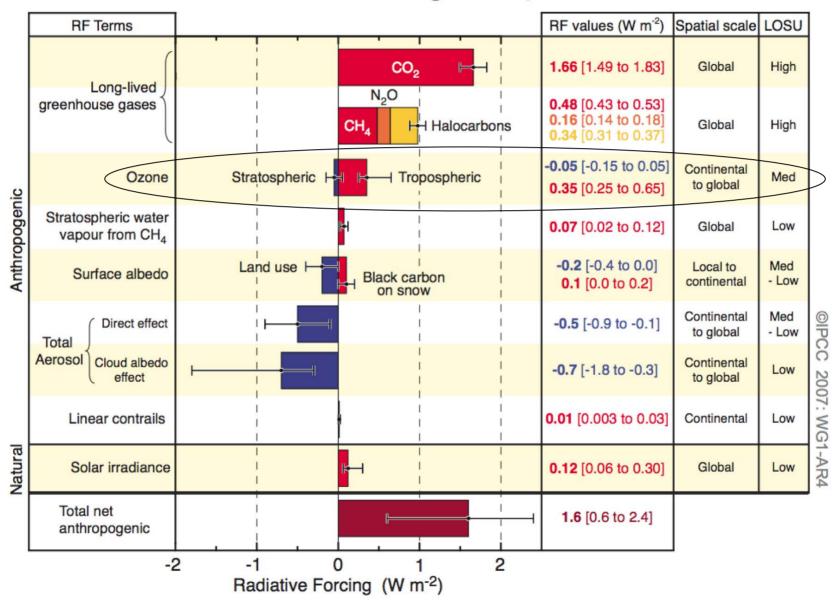
- TES retrieves ozone profiles from spectrally resolved IR radiances.
- Using both spectra and retrievals, we can examine the sources of variability in the outgoing flux for the IR ozone absorption band.

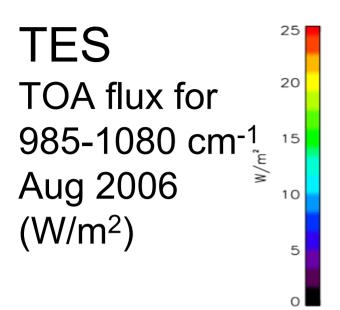


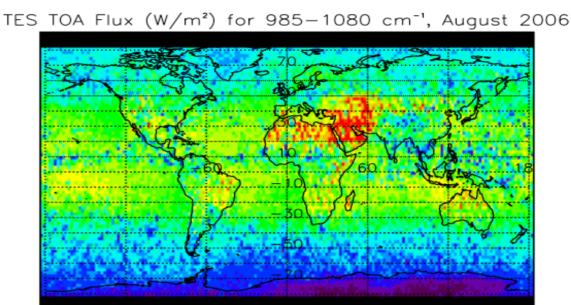


#### 4th IPCC Assessment (2007)

#### **Radiative Forcing Components**

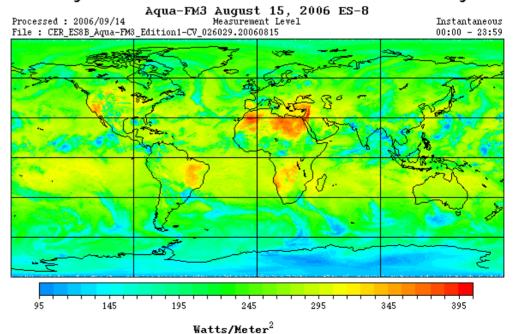




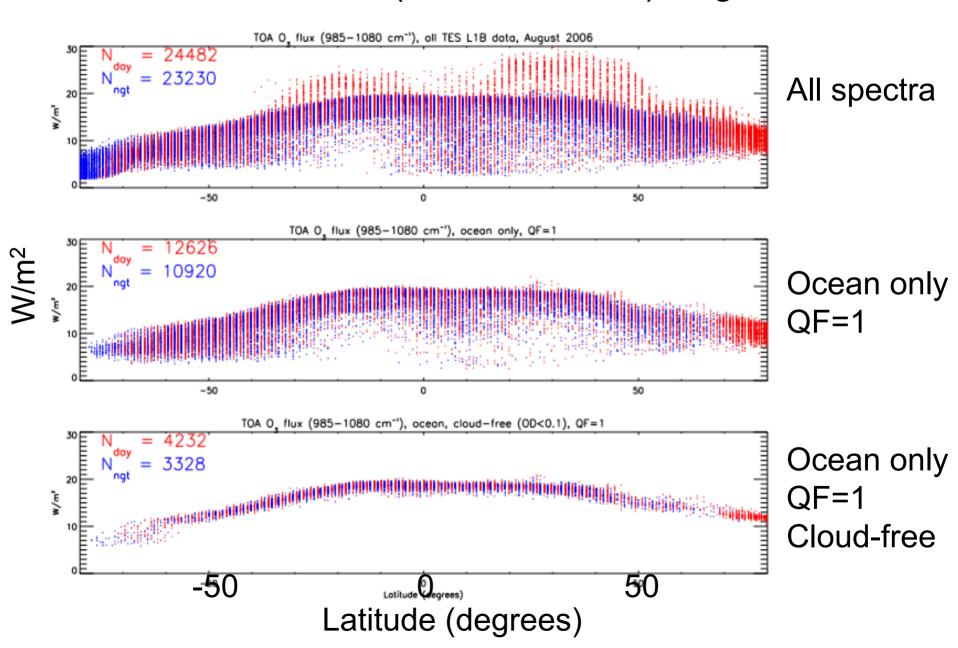


# CERES Longwave TOA flux Aug 15, 2006 (W/m²)

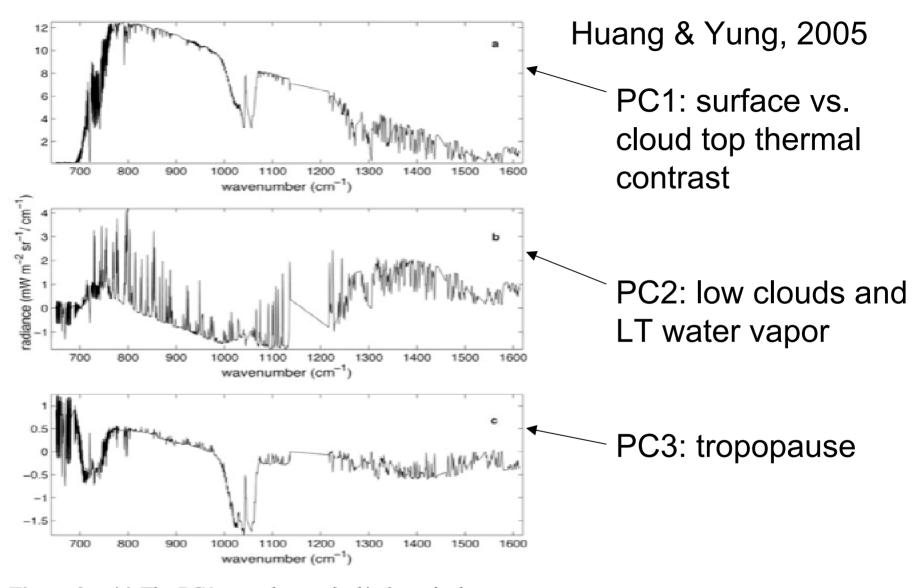




#### TES TOA flux (985-1080 cm<sup>-1</sup>) Aug 2006



#### AIRS nadir spectra principal component analysis



**Figure 3.** (a) The PC1 over the tropical/subtropical oceans between 32°S and 32°N derived from AIRS spectra collected during July 1–16, 2003. (b) The PC2. (c) The PC3.

#### AIRS nadir spectra principal component analysis

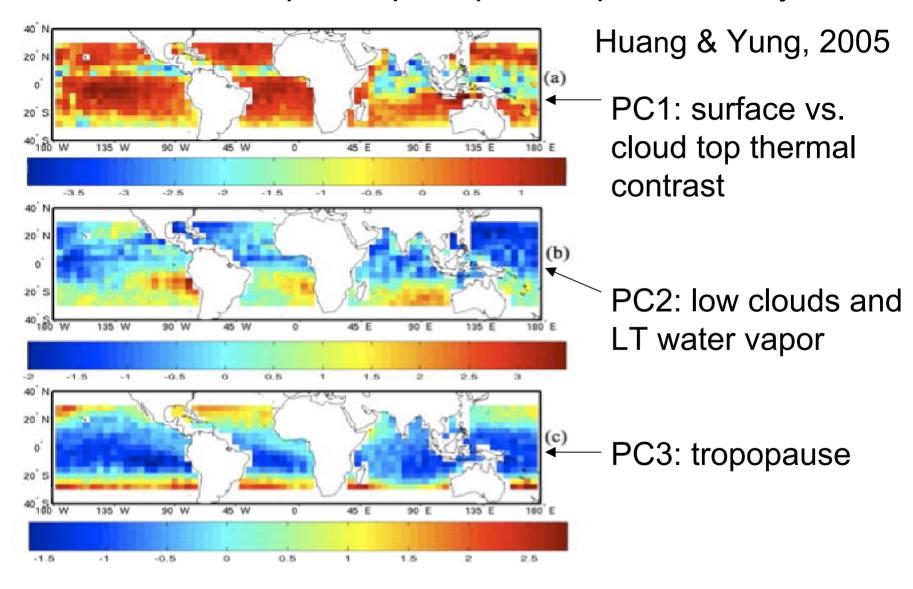


Figure 4. (a) The spatial map of the PC1 shown in Figure 3a. In other words, it is the normalized expansion coefficient of the PC1 computed based on equation (3) in the context. (b) The spatial map of the PC2 shown in Figure 3b. (c) The spatial map of the PC3 shown in Figure 3c.

#### SVD for cloud-free, ocean, tropics TES spectra

Average radiance JJA 2006 tropics 985-1080 cm<sup>-1</sup>

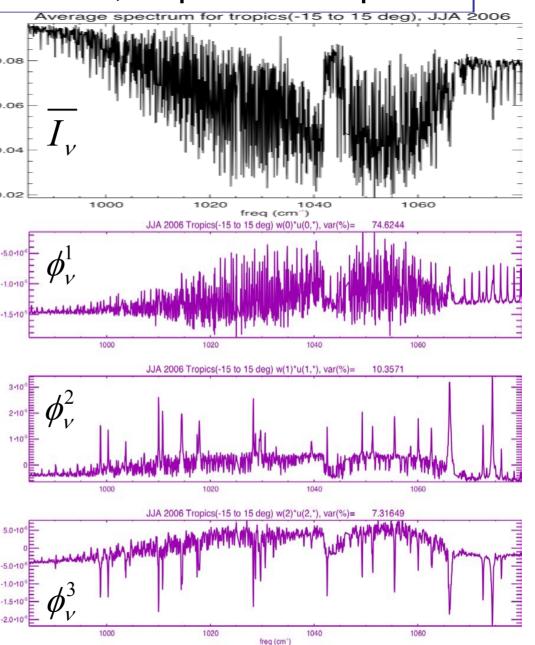
1<sup>st</sup> 3 Singular Vectors (93.7% of variance)

Primary variability from:

 $\phi^1_{\nu}$  SST

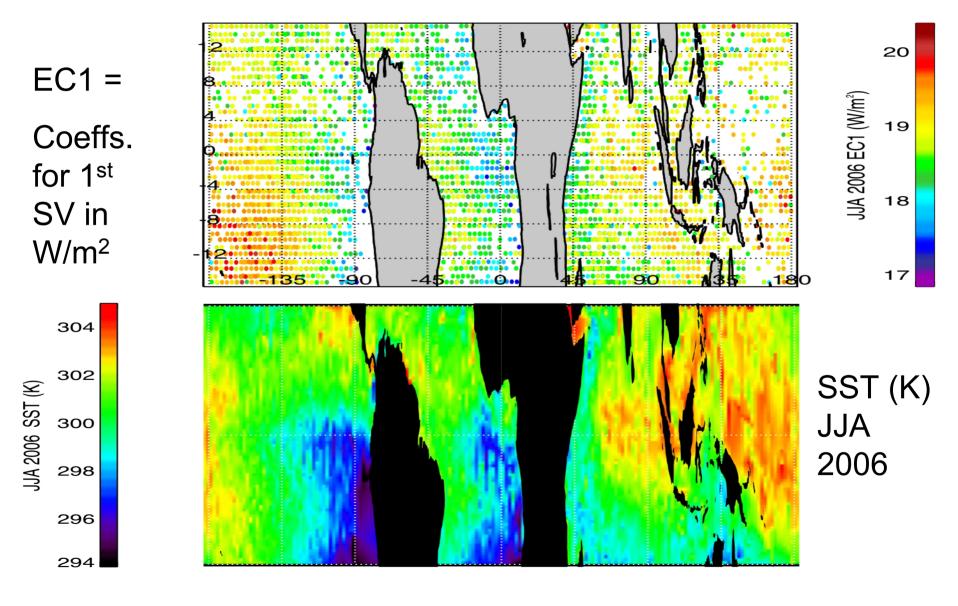
 $\phi_{\nu}^{2}$  water vapor

 $\phi_{\nu}^{3}$  upper trop. O3

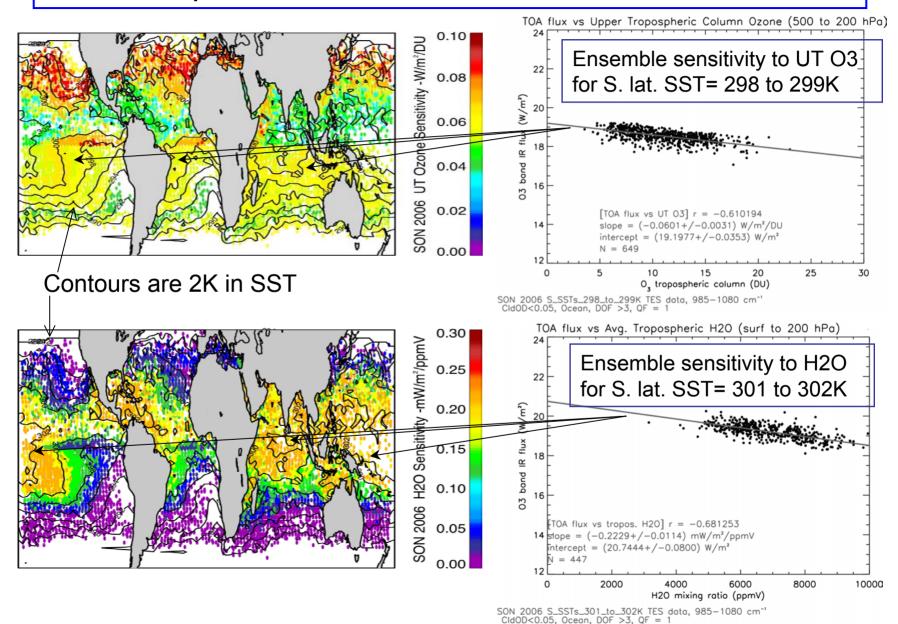


## Maps of integrated expansion coefficients:

$$EC_{i}(x) = \int_{V} \left[ \left\langle \left[ I_{v}(x) - \overline{I}_{v} \right], \phi_{v}^{(i)} \right\rangle \phi_{v}^{(i)} + \overline{I}_{v} \right] dv$$

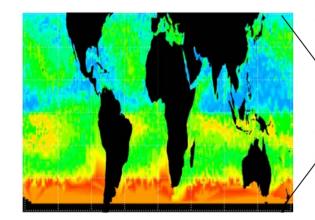


## Linear fits of TOA flux (W/m<sup>2</sup>) vs ozone and water vapor N/S hemisphere SST bin ensembles, SON 2006



## Normalized Radiative Forcing (W/m²/DU)

TES JJA 2006 ensemble sensitivities



TES global, annual (45 \$\pi\$ to 45 \$\pi\$ N) TOA avg.

 $= 0.055 \text{ W/m}^2/\text{DU}$ (0.017 st.dev.)

Model range

JJA 2006 UT Ozone Sensitivity -W/m²/DU

0.08

0.06

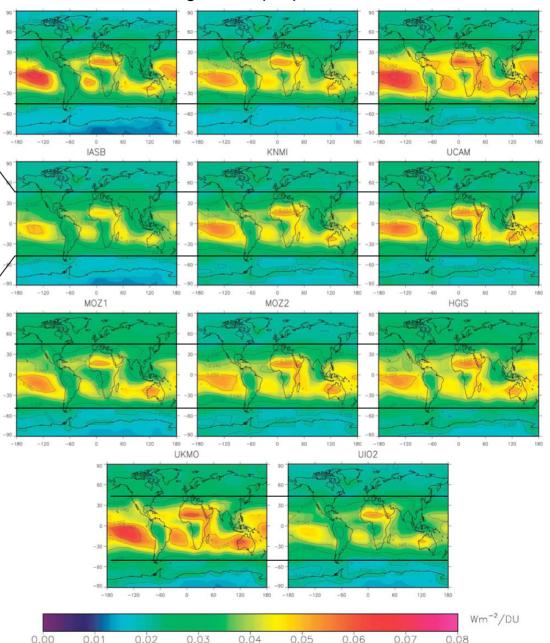
0.04

 $= 0.042 - 0.052 \text{ W/m}^2/\text{DU}$ 

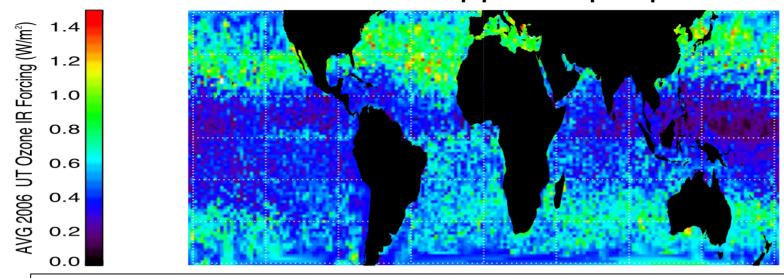
LW clear, inst., all lats,

trop only (no strat)

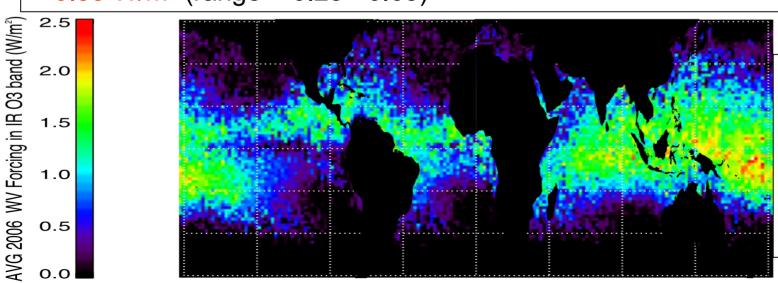
Gauss et al., JGR, 2003, Radiative Forcing in 21st Century in 17 due to ozone changes in troposphere/lower strat.



#### OLR reduction due to upper tropospheric ozone

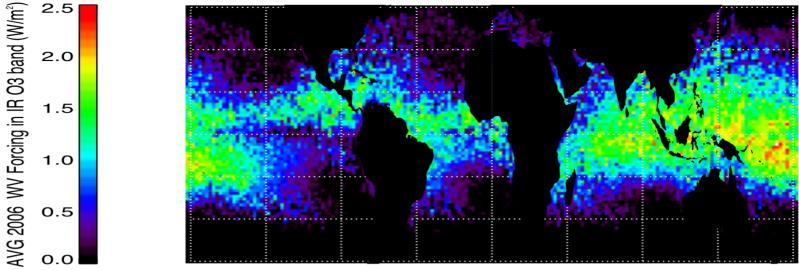


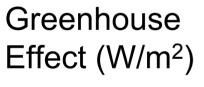
TES global, annual avg =  $0.48 \text{ W/m}^2$  (0.24 std) for  $45 \oplus \text{S}$  to  $45 \oplus \text{N}$  IPCC (2007) value for anthropogenic tropospheric ozone radiative forcing =  $0.35 \text{ W/m}^2$  (range = 0.25 - 0.65)



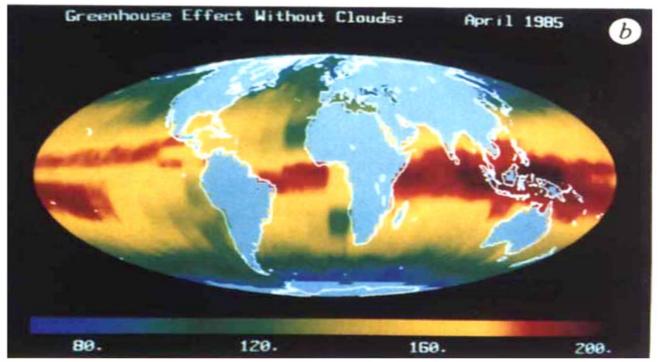
Reduced
OLR from
water vapor
absorption
in IR ozone
band

Reduced OLR from water vapor in IR ozone band



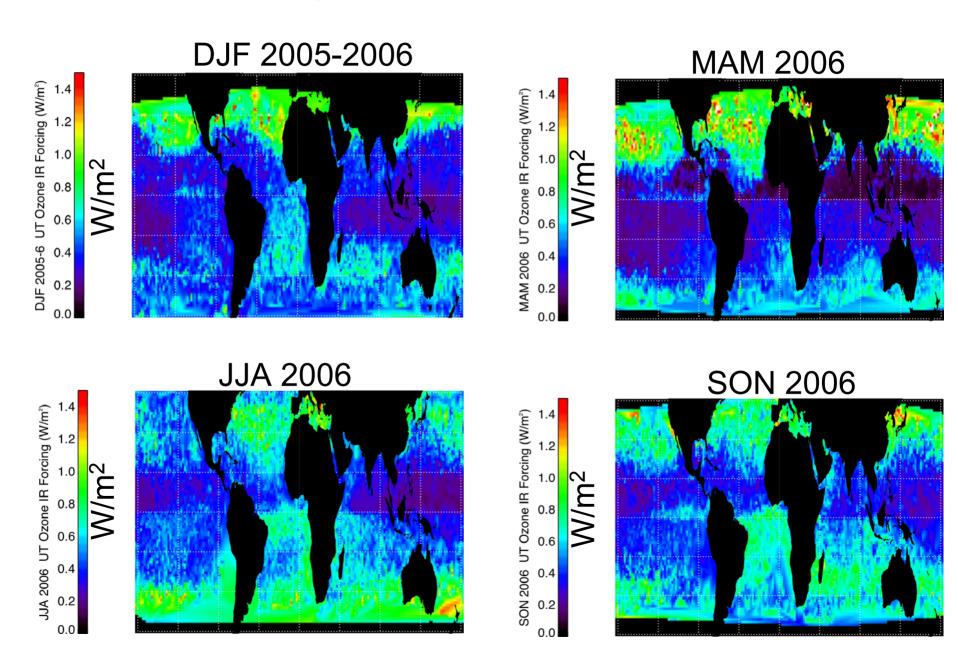


$$G = \sigma T^4 - F_{TOA}$$

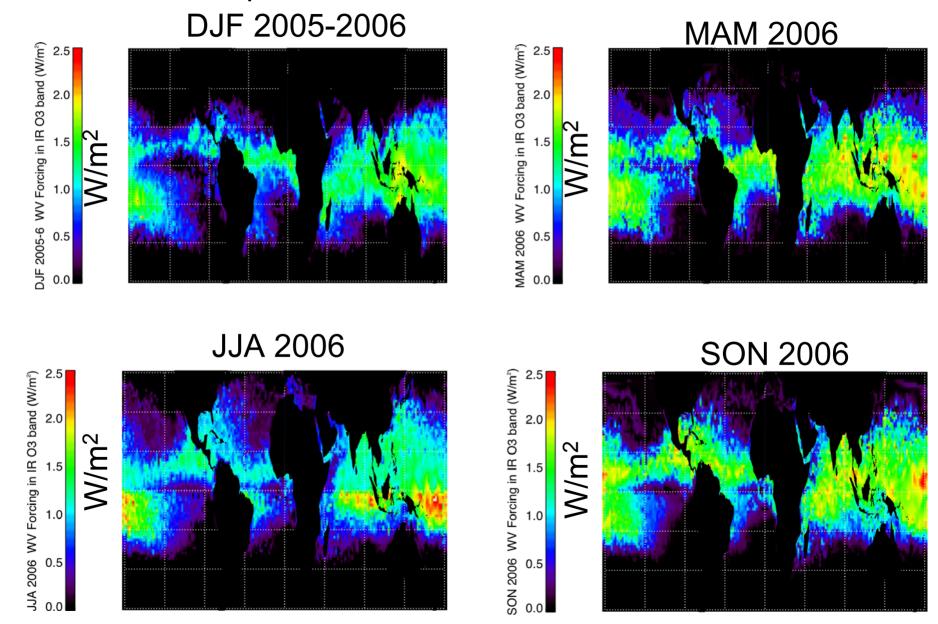


A. Ravel & V. Ramanathan, Nature, 1989

#### Seasonal dependence of UT ozone OLR reduction



Seasonal dependence of OLR reduction due to water vapor in IR ozone band

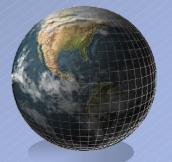


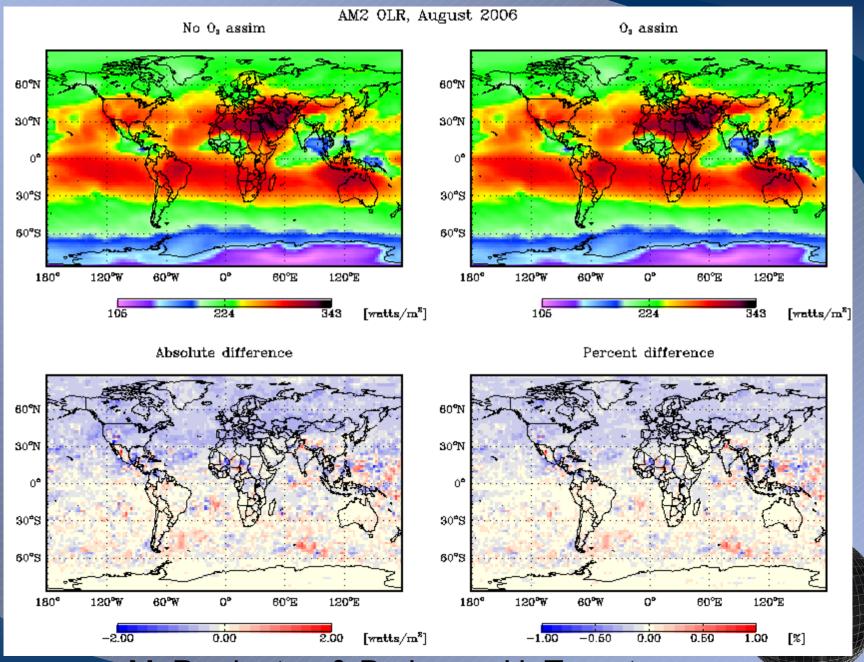
#### Conclusions for OLR sensitivity to tropospheric ozone

- Using TES spectra, retrieved surface & cloud properties and atmospheric profiles, we can investigate the processes that drive OLR variability.
- We estimate OLR sensitivity to ozone and water vapor by constructing ensemble observations binned by SST. We find an annual average OLR sensitivity to upper trop. ozone of 0.055 W/m²/DU (stdev = 0.017). This is comparable to model estimates but with more sensitivity in the northern hemisphere.
- Using estimated sensitivities, we find the average OLR reduction = 0.48 W/m² (stdev = 0.24 std) for upper trop. ozone (global, annual avg., -45⊕ to 45⊕)
  - IPCC (2007) value =  $0.35 \text{ W/m}^2$  (range = 0.25 0.65) for anthropogenic trop. ozone

#### Future directions

- Using TES observations of OLR and ozone to test the processes that drive OLR variability in climate models.
- Investigate OLR sensitivity for other atmospheric species e.g. H2O, CH4, CO2 for any observation altitude/lat/lon.
   -based on computed jacobians
- Characterize OLR sensitivity for different surface & cloud conditions
- Separate anthropogenic and natural contributions using inverse modeling



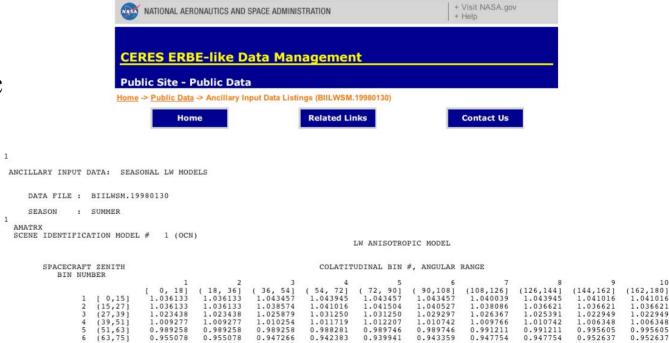


M. Parrington & D. Jones, U. Toronto

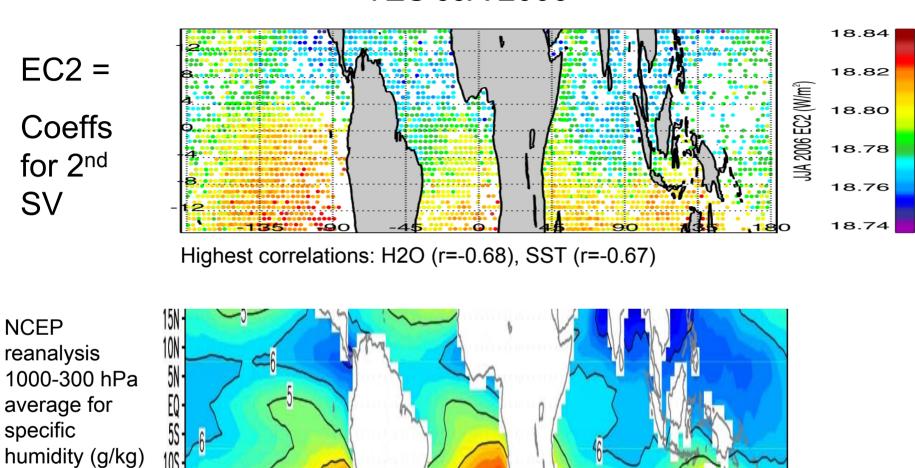
#### Radiance-to-Flux Conversion

 $M = \pi L / R(\theta)$  With M = flux estimate in W/m<sup>2</sup>  $L = \text{radiance in W/m}^2/\text{sr}$   $R(\theta) = \text{anisotropy from an angular dist. model (ADM)}.$ For LW, R depends on the viewing angle,  $\theta$ .
(from ERBE references)
For now, using R=1.05 for nadir ( $\theta = 0$ )

Could use season/latitude/srfc values for R from CERES - range ~ 1.03 to 1.06



#### **TES JJA 2006**



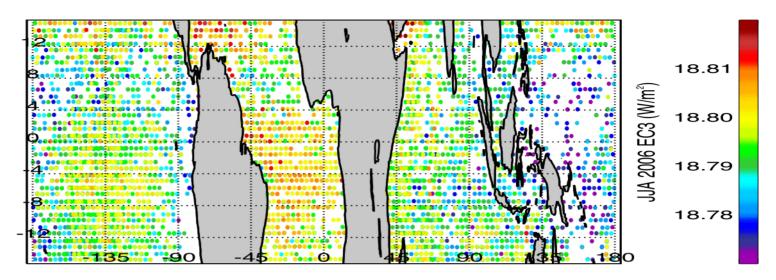


NOAA/ESRL Physical Sciences Division

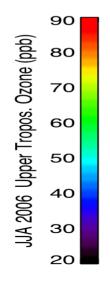
MAX=8.21216 MIN=1.86098 NCEP GrADS image

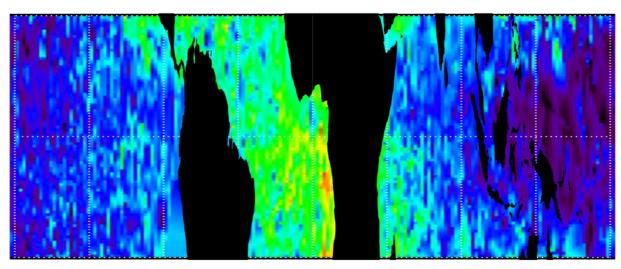
#### **TES JJA 2006**

EC3 =
Coeffs
for 3<sup>rd</sup>
SV



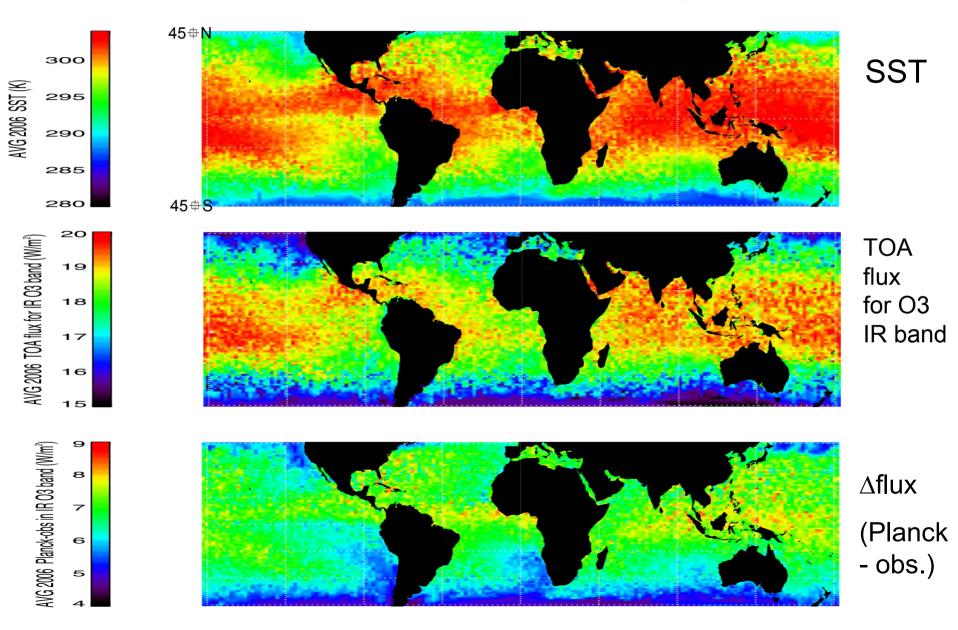
Highest correlation: UT O3 (r=0.60)



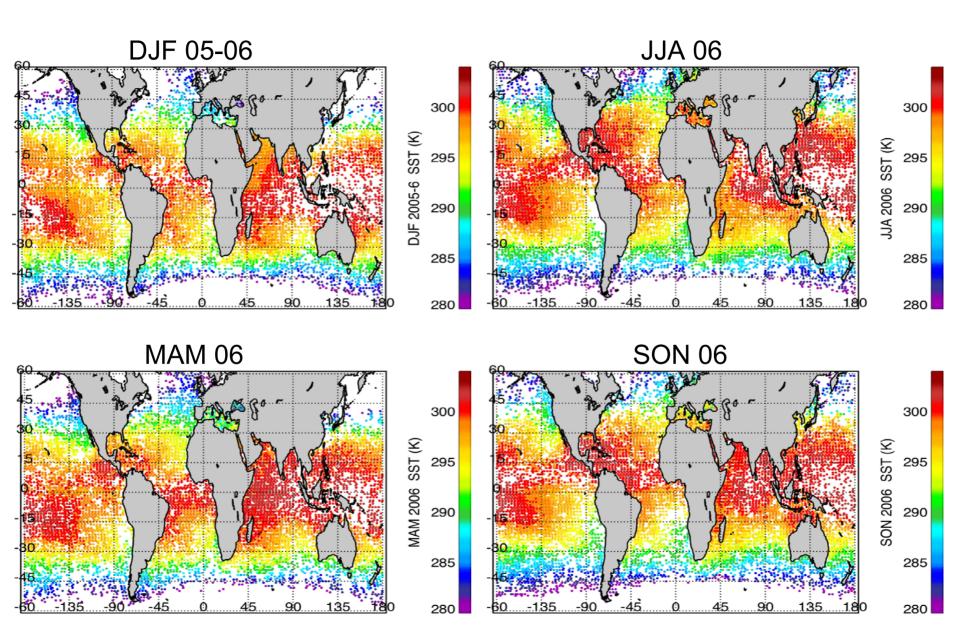


UT O3 (ppb) JJA 2006 (500-200 hPa)

#### TES 2006 annual averages

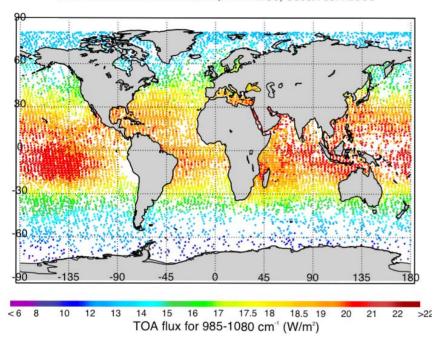


#### Sea Surface Temperature (K) - cloud-free ocean scenes



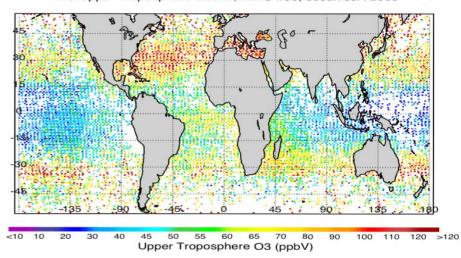
#### TOA flux and ozone for JJA 2006

TES TOA flux for ozone band, cloud-free, ocean JJA 2006

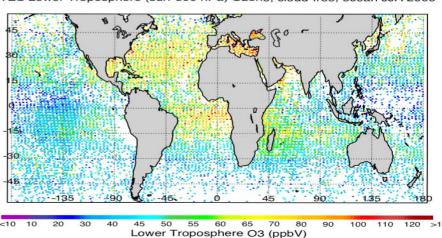


UT = 500-200 hPa for |lat| < 35° = 500-300 hPa for 35°-55°

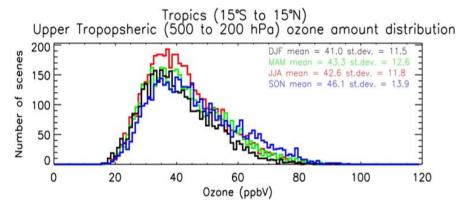
TES Upper Troposphere Ozone, cloud-free, ocean JJA 2006

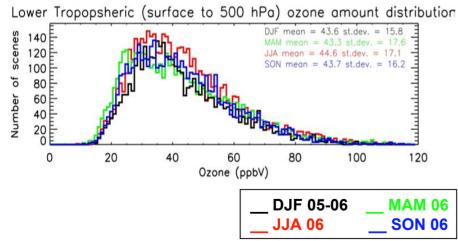


TES Lower Troposphere (surf-500 hPa) Ozone, cloud-free, ocean JJA 2006

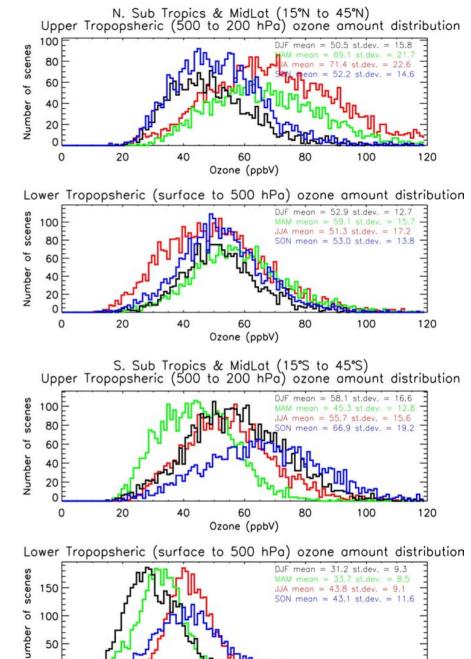


# Upper & Lower Troposphere O3 distributions by season, Tropics, N., S.





Only clear-sky, ocean profiles are included.



80

Ozone (ppbV)

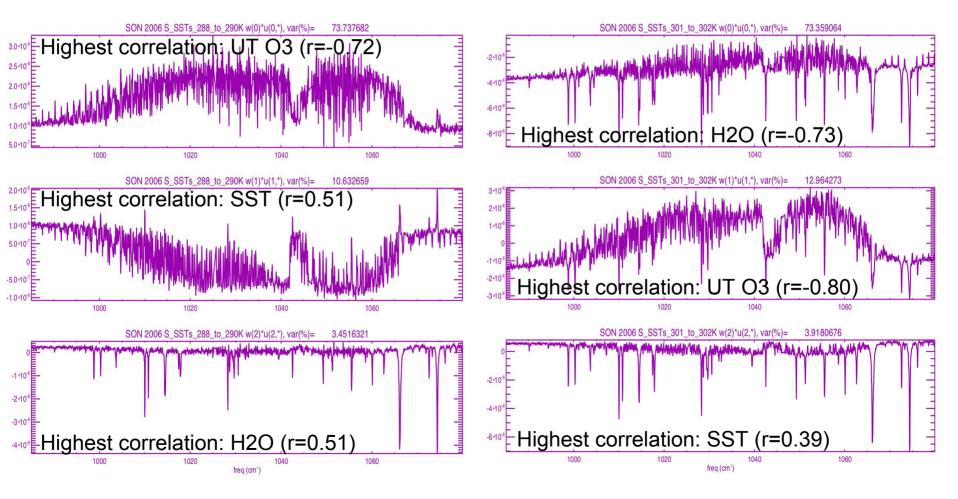
100

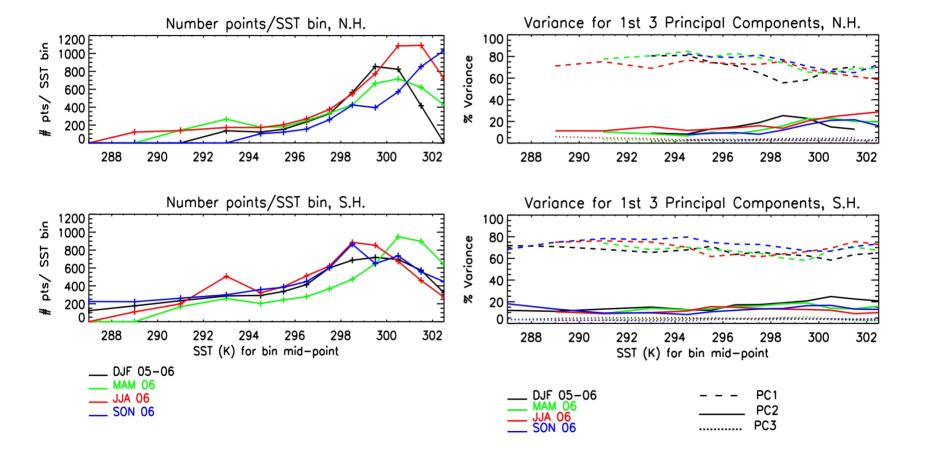
120

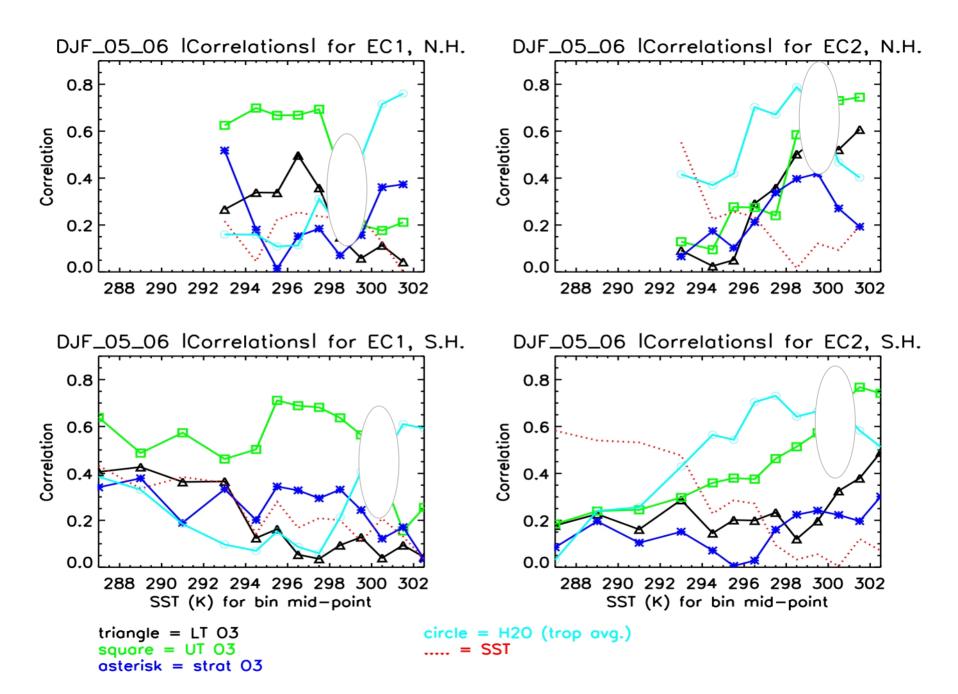
20

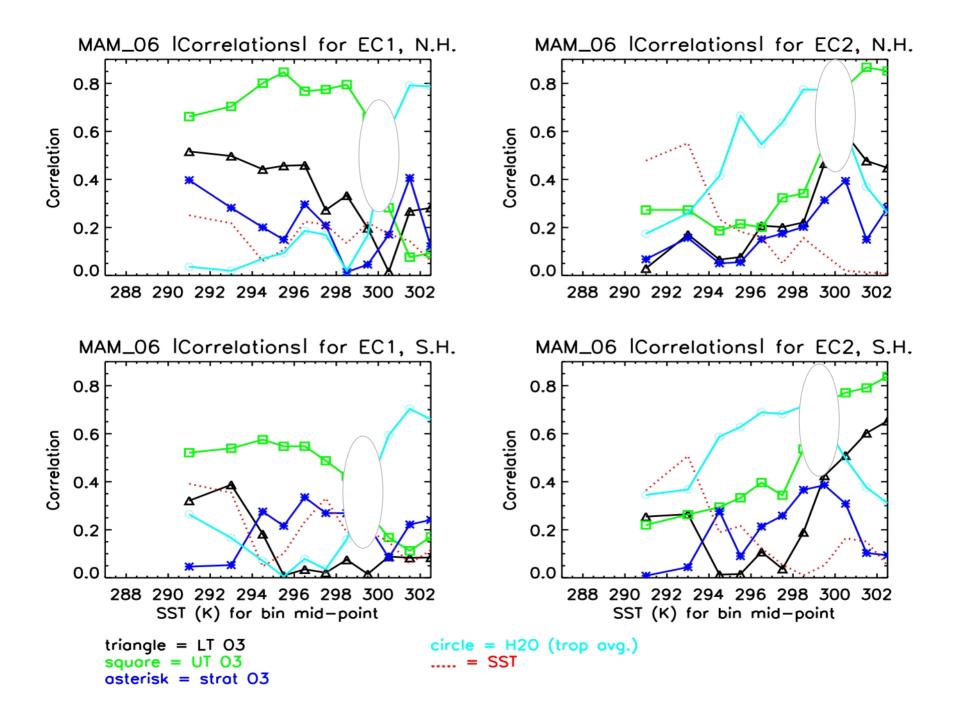
Bin TES measurements by Sea Surface Temperature (SST) to isolate variability in TOA flux due to Ozone and H2O

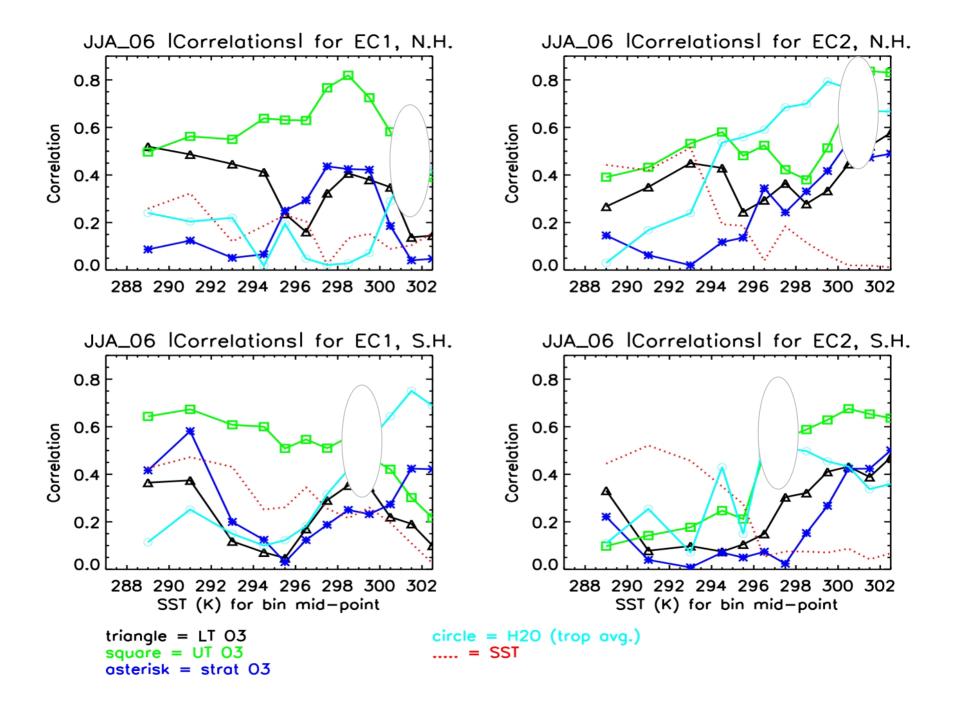
1<sup>st</sup> 3 Singular Vectors for SON 2006, cloud-free, ocean, SST bins S. midlat [288K to 290K] and S. tropics [301K to 302K]

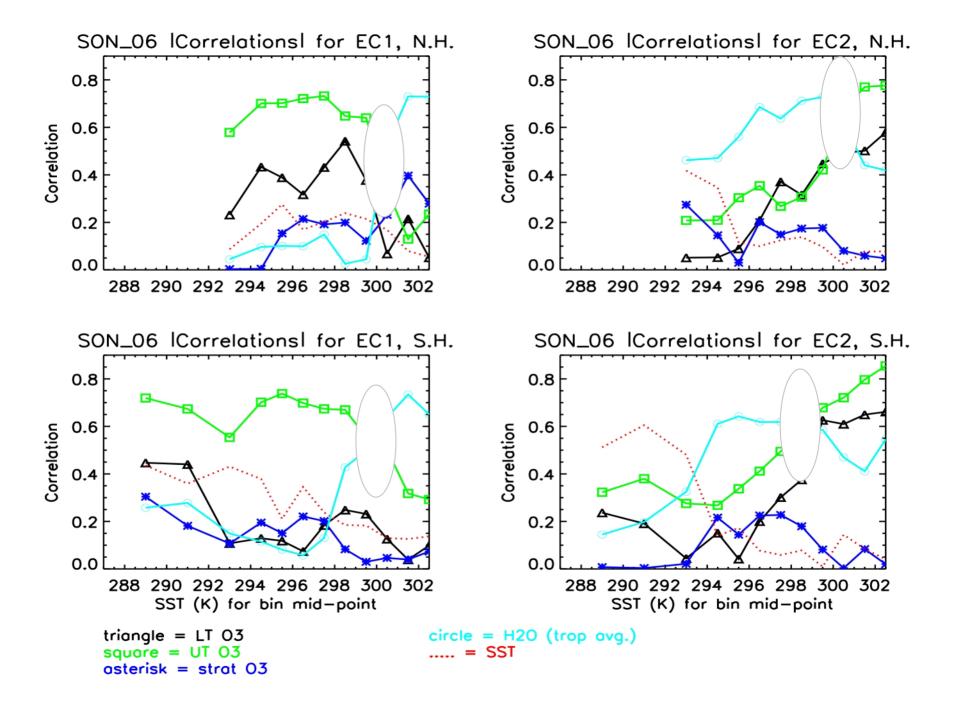


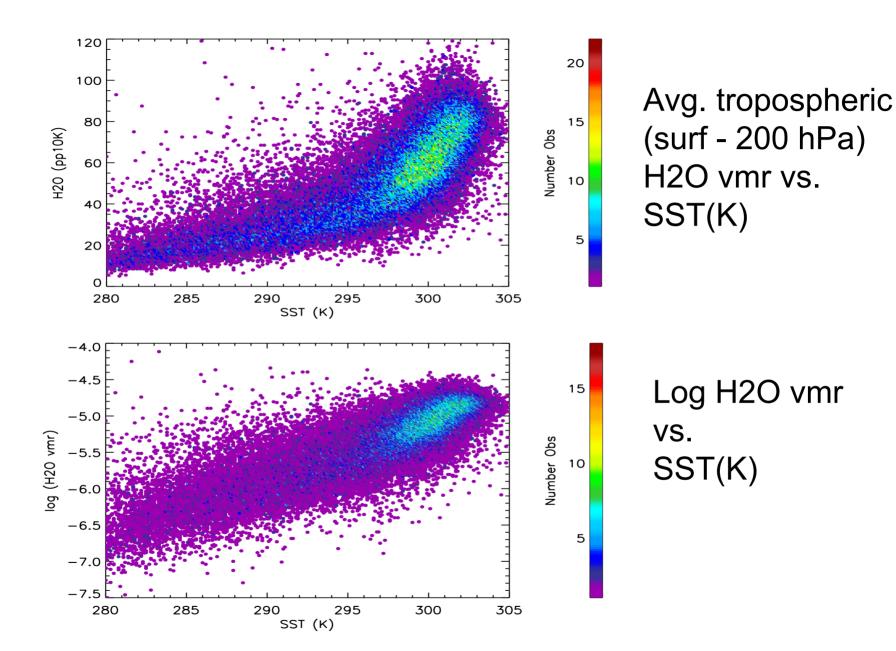


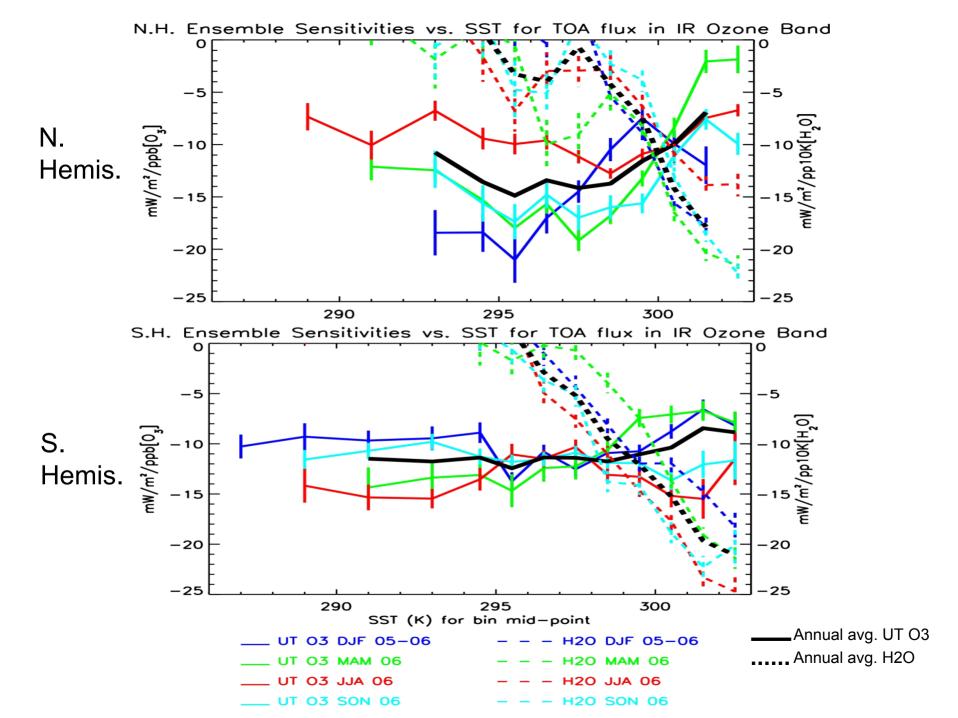












## Linear fits of TOA flux (W/m²) vs O<sub>3</sub> (pppV) for SST bin, N/S hemisphere ensembles, JJA\_06

